

Approximating the complex-valued Green-tensor amplitude by a real-valued Green-tensor amplitude

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Summary

The paper is devoted to the approximation of the complex-valued Green-tensor amplitude by a real-valued Green-tensor amplitude with a phase shift. This approximation is required in the method of Jan Šílený for determining the real-valued seismic moment tensor from the maximum real-valued vectorial amplitude picked in the polarization diagram.

Keywords

Wave propagation, elastic waves, Green tensor, amplitude, seismic moment tensor.

1. Introduction

The tensorial amplitude of the S-wave Green tensor is often complex-valued, especially if the wavefield is recorded at the Earth surface (Červený, 2001).

The method of Šílený & Milev (2008), Šílený et al. (2009) and Horálek & Šílený (2013) for determining the real-valued seismic moment tensor from the maximum real-valued vectorial amplitude picked in the polarization diagram requires the approximation of the complex-valued Green-tensor amplitude by a real-valued Green-tensor amplitude with a phase shift. When studying this approximation, we shall assume the ray-theory approximation of the Green tensor expressed in terms of the 3×3 tensorial amplitude and travel time. The approximation is applicable to Green-tensor amplitudes in both isotropic and anisotropic heterogeneous media, including the isotropic ray theory, the anisotropic ray theory, and the prevailing-frequency approximation of the coupling ray theory for S waves (Klimeš & Bulant, 2012).

2. Approximation of the complex-valued Green-tensor amplitude

Complex-valued Green-tensor amplitude \mathbf{G}^C maps seismic force \mathbf{f} onto the complex-valued displacement \mathbf{u}^C ,

$$\mathbf{u}^C = \mathbf{G}^C \mathbf{f} \quad . \quad (1)$$

We wish to approximate complex-valued Green-tensor amplitude \mathbf{G}^C by real-valued Green-tensor amplitude \mathbf{G}^R with phase shift factor $\exp(i\varphi)$, which map seismic force \mathbf{f} onto the approximate complex-valued displacement \mathbf{u}^R ,

$$\mathbf{u}^R = \mathbf{G}^R \mathbf{f} \exp(i\varphi) \quad . \quad (2)$$

The square of the wavefield error of this approximation is

$$\delta^2 = (\mathbf{u}^C - \mathbf{u}^R) + (\mathbf{u}^C - \mathbf{u}^R) \quad , \quad (3)$$

where $^+$ denotes the Hermitian adjoint. We insert relations (1) and (2) into definition (3) and obtain

$$\delta^2 = \mathbf{f}^+ [\mathbf{G}^C - \mathbf{G}^R \exp(i\varphi)]^+ [\mathbf{G}^C - \mathbf{G}^R \exp(i\varphi)] \mathbf{f} \quad . \quad (4)$$

We average the square of wavefield error δ^2 over all spatial directions of unit seismic force \mathbf{f} ,

$$\langle \delta^2 \rangle = \frac{1}{3} \text{Tr}\{ [\mathbf{G}^C - \mathbf{G}^R \exp(i\varphi)]^+ [\mathbf{G}^C - \mathbf{G}^R \exp(i\varphi)] \} \quad . \quad (5)$$

After multiplication, relation (5) reads

$$\langle \delta^2 \rangle = \frac{1}{3} \text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C - (\mathbf{G}^R)^T \mathbf{G}^C \exp(-i\varphi) - (\mathbf{G}^C)^+ \mathbf{G}^R \exp(i\varphi) + (\mathbf{G}^R)^T \mathbf{G}^R] \quad , \quad (6)$$

where T denotes the transpose. We decompose the complex-valued Green-tensor amplitude into its real and imaginary parts and obtain

$$\langle \delta^2 \rangle = \frac{1}{3} \text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C - 2(\mathbf{G}^R)^T \text{Re}(\mathbf{G}^C) \cos(\varphi) - 2(\mathbf{G}^R)^T \text{Im}(\mathbf{G}^C) \sin(\varphi) + (\mathbf{G}^R)^T \mathbf{G}^R] \quad . \quad (7)$$

3. Minimizing the wavefield error

We wish to select real-valued Green-tensor amplitude \mathbf{G}^R and phase shift φ so that the average square (7) of wavefield error is minimal.

Differentiating wavefield error (7) with respect to real-valued Green-tensor amplitude \mathbf{G}^R , we obtain equation

$$\frac{2}{3} [-\text{Re}(\mathbf{G}^C) \cos(\varphi) - \text{Im}(\mathbf{G}^C) \sin(\varphi) + \mathbf{G}^R] = \mathbf{0} \quad (8)$$

for the minimum of the average square (7) of wavefield error.

Differentiating wavefield error (7) with respect to phase shift φ , we obtain equation

$$\frac{2}{3} \text{Tr}[(\mathbf{G}^R)^T \text{Re}(\mathbf{G}^C) \sin(\varphi) - (\mathbf{G}^R)^T \text{Im}(\mathbf{G}^C) \cos(\varphi)] = 0 \quad . \quad (9)$$

for the minimum of the average square (7) of wavefield error.

Equation (8) yields

$$\mathbf{G}^R = \text{Re}(\mathbf{G}^C) \cos(\varphi) + \text{Im}(\mathbf{G}^C) \sin(\varphi) \quad . \quad (10)$$

We insert relation (10) into equation (9) and arrive at

$$\begin{aligned} \frac{2}{3} \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C) \cos(\varphi) \sin(\varphi) + \text{Im}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C) \sin(\varphi) \sin(\varphi) \\ - \text{Re}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C) \cos(\varphi) \cos(\varphi) - \text{Im}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C) \sin(\varphi) \cos(\varphi)] = 0 \quad , \quad (11) \end{aligned}$$

which yields

$$\text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C) - \text{Im}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] \sin(2\varphi) = 2 \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] \cos(2\varphi) \quad . \quad (12)$$

The solution of equation (12) is

$$\tan(2\varphi) = 2 \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] / \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C) - \text{Im}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] \quad . \quad (13)$$

We can thus calculate phase shift φ using relation (13) and insert it into relation (10) to obtain the best real-valued Green-tensor amplitude \mathbf{G}^R .

4. Relative wavefield error

For the best real-valued Green-tensor amplitude (10), the average square (7) of wavefield error reads

$$\langle \delta^2 \rangle = \frac{1}{3} \text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C - (\mathbf{G}^R)^T \mathbf{G}^R] \quad . \quad (14)$$

The square of the norm of the complex-valued displacement is

$$|\mathbf{u}^C|^2 = \mathbf{f}^+ (\mathbf{G}^C)^+ \mathbf{G}^C \mathbf{f} \quad . \quad (15)$$

We average the square of the norm over all spatial directions of unit seismic force \mathbf{f} ,

$$\langle |\mathbf{u}^C|^2 \rangle = \frac{1}{3} \text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C] \quad . \quad (16)$$

The relative root mean square error of the approximation of the complex-valued Green-tensor amplitude by the real-valued Green-tensor amplitude is

$$\rho = \sqrt{\langle \delta^2 \rangle / \langle |\mathbf{u}^C|^2 \rangle} \quad , \quad (17)$$

where $\langle \delta^2 \rangle$ and $\langle |\mathbf{u}^C|^2 \rangle$ are given by expressions (14) and (16) with

$$\text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C] = \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C)] + \text{Tr}[\text{Im}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] \quad . \quad (18)$$

If relative error (17) is acceptable, the real-valued Green-tensor amplitude can be used in the inversion according to Šílený & Milev (2008), Šílený et al. (2009), and Horálek & Šílený (2013). If relative error (17) is not acceptable, the corresponding data from the receiver should not be used in that inversion.

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